

# On a dynamical separation of waves and balanced motions based on surface measurements

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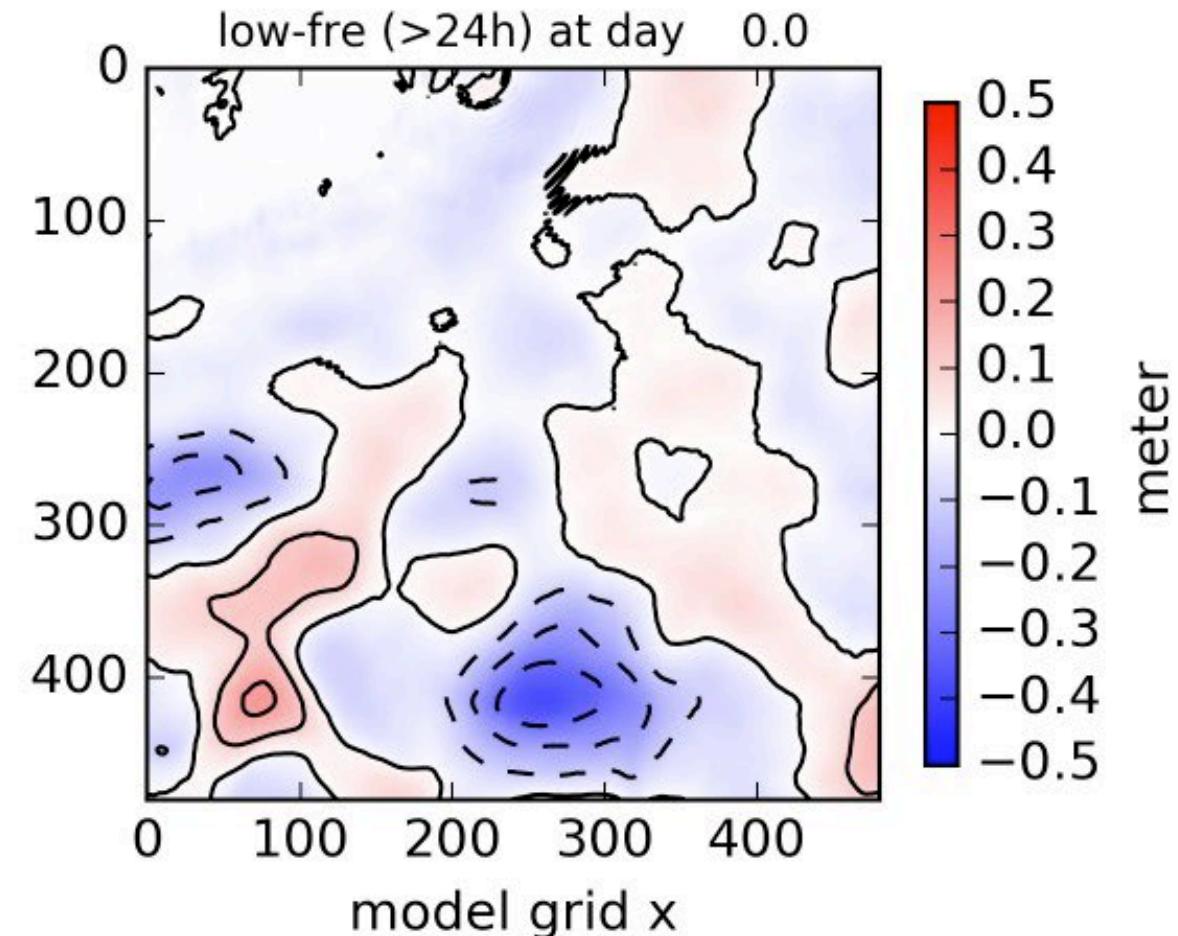
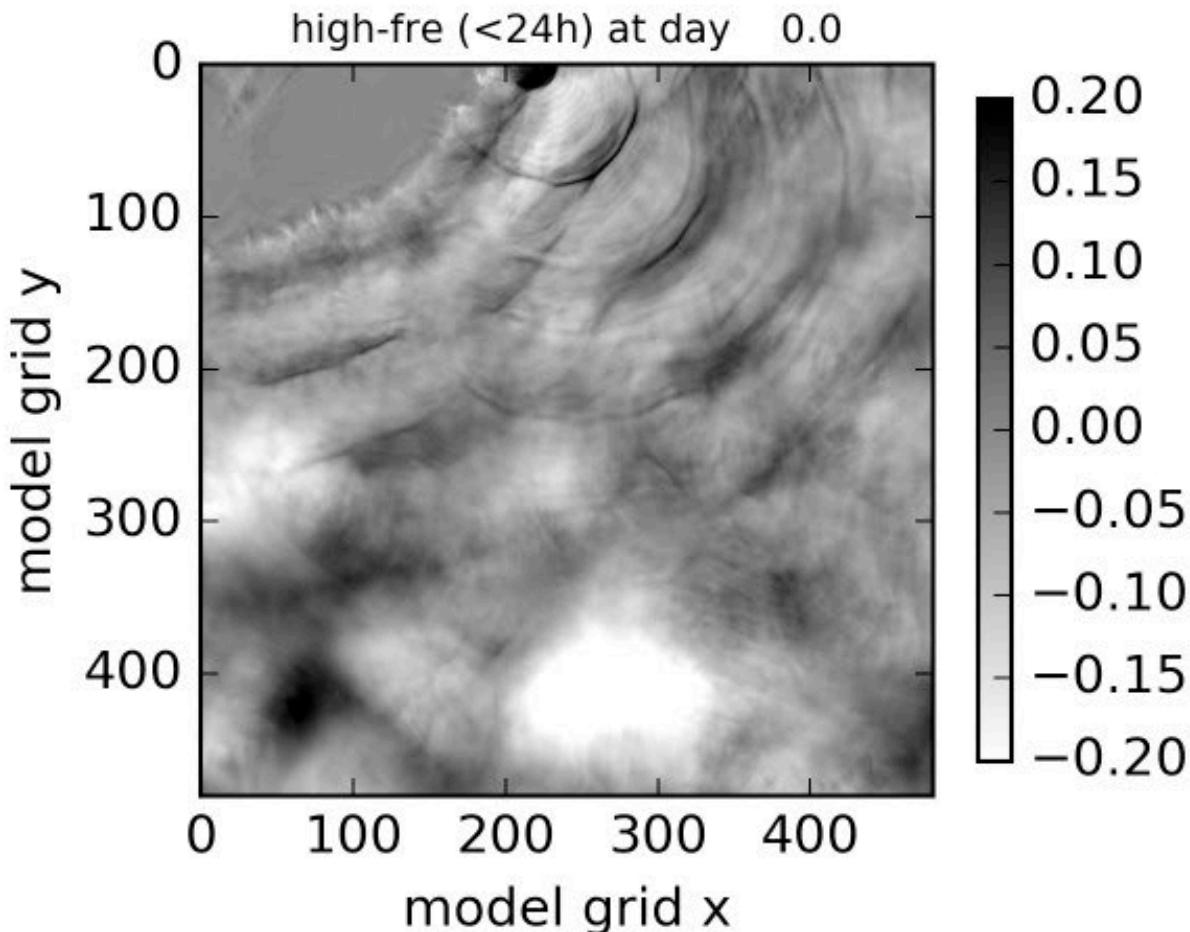
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SWOT ST meeting @Toulouse, France

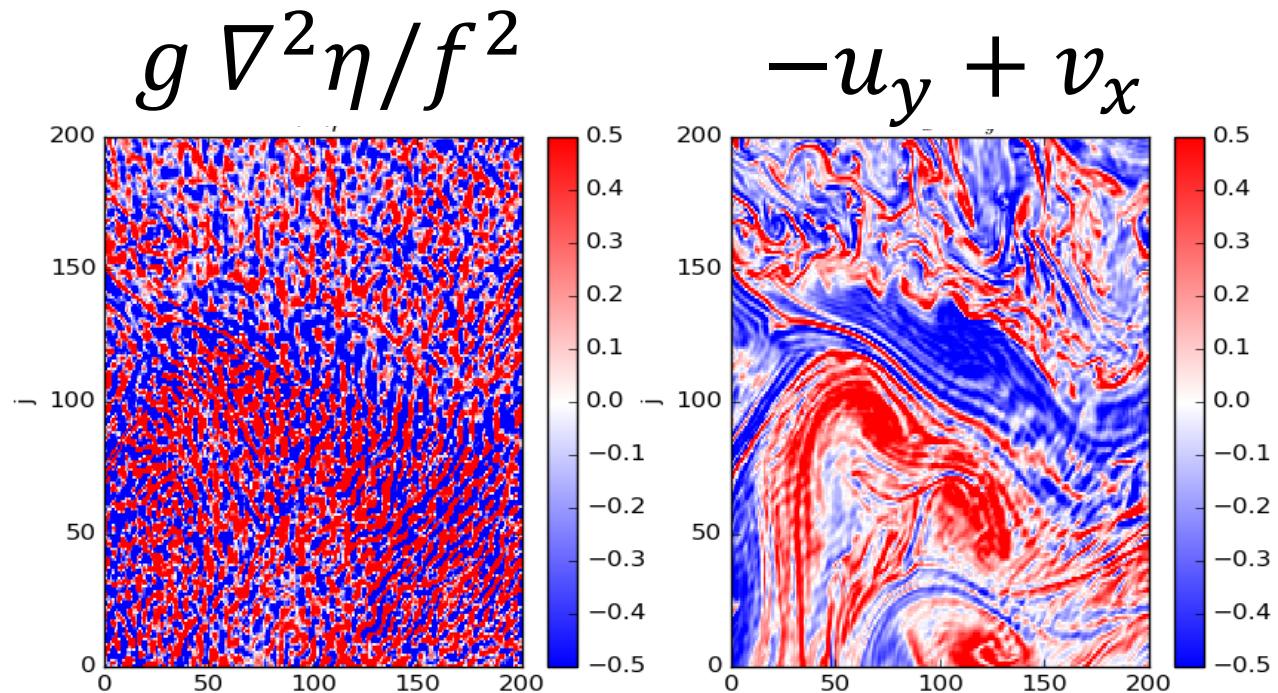
6/27/2017

# The coexistence of the waves and balanced motions in SSH



SSH from 1/48 MITgcm simulation

The simple Laplacian of  
SSH is not relative  
vorticity.



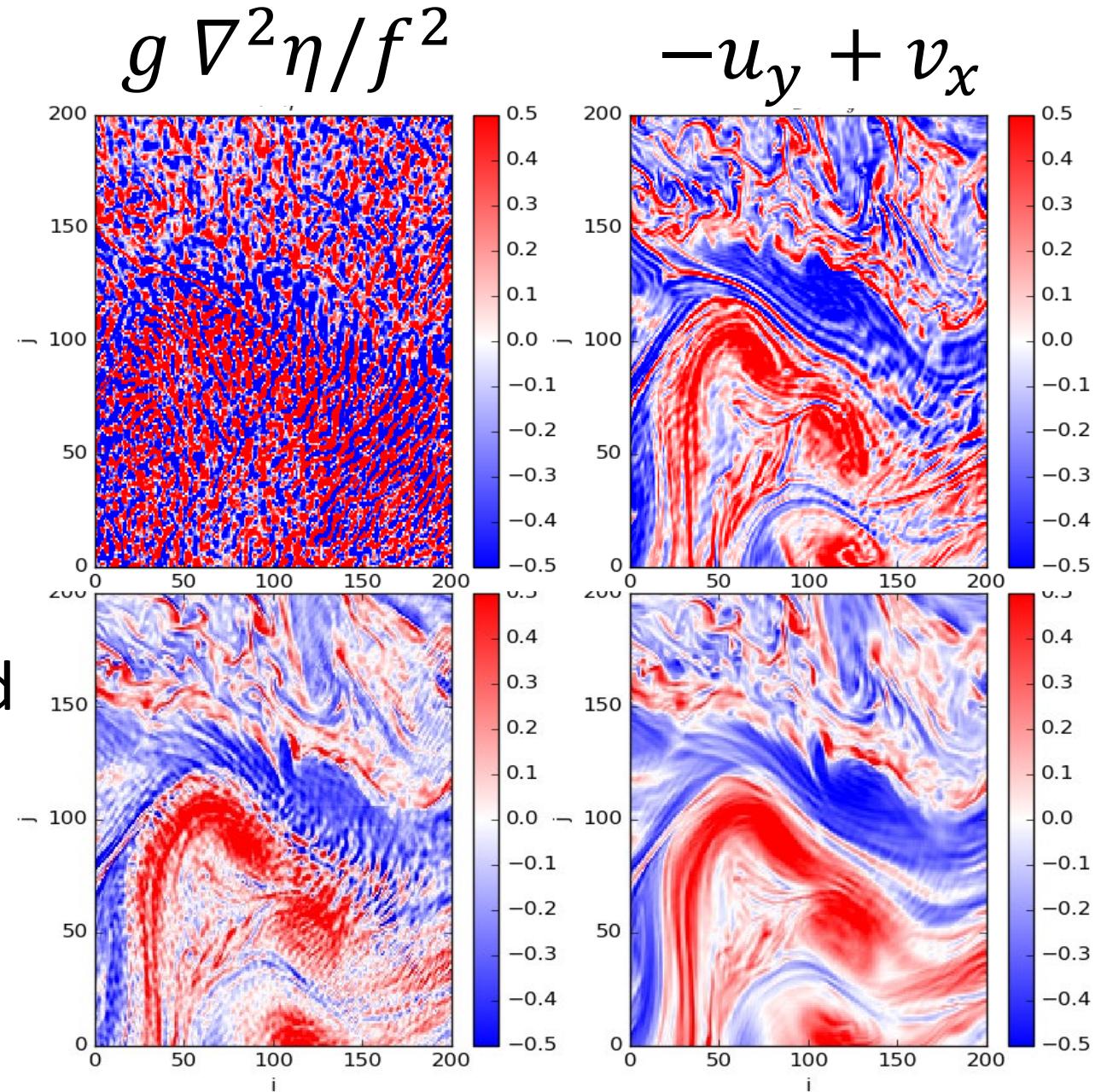
Snapshot

Q: How to separate internal waves and balanced motions?

A: Daily average.

The simple Laplacian of SSH is not relative vorticity.

Daily average does yield results approaching balanced dynamics.



Snapshot  
Daily average

Can we regenerate the SSH time evolution  
based on satellite snapshots?

What are the requirements?

We need an initial condition with

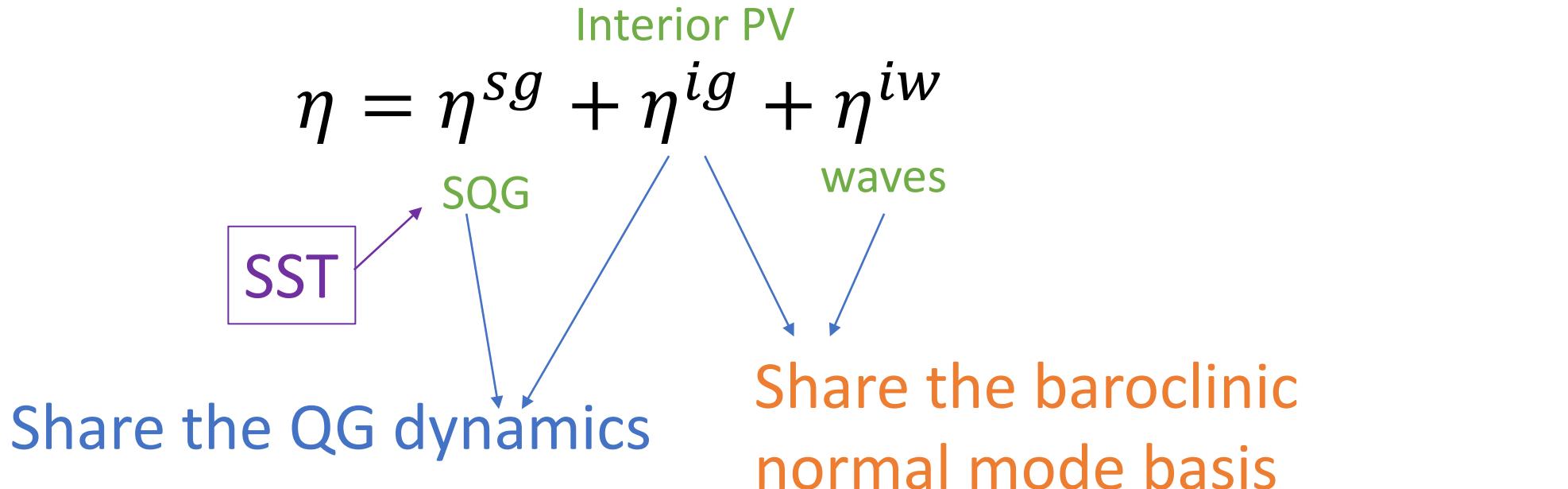
$$\eta(x, y)$$

$$\rho(x, y, z)$$

$$\vec{u}_H(x, y, z)$$

# interior+surface Quasi-geostrophy+Waves

Based on Wang et al. (2013) and Ponte et al. (2017)



Wang et al. (2013):  $\eta^{ig} = \eta - \eta^{sg}$

Does not distinguish  $\eta^{ig}$  and  $\eta^{iw}$

Ponte et al. (2017):

$$\eta^{iw} = \eta - \eta^{sg}(eSQG)$$

Miss the interior PV not resolved by eSQG

# An idealized demo the control run

$$\omega = 1.41 \times 10^{-4} s^{-1}$$

$$f = 8 \times 10^{-5} s^{-1}$$

$$H = -5000 m$$

$$dx = dy = 2 km$$

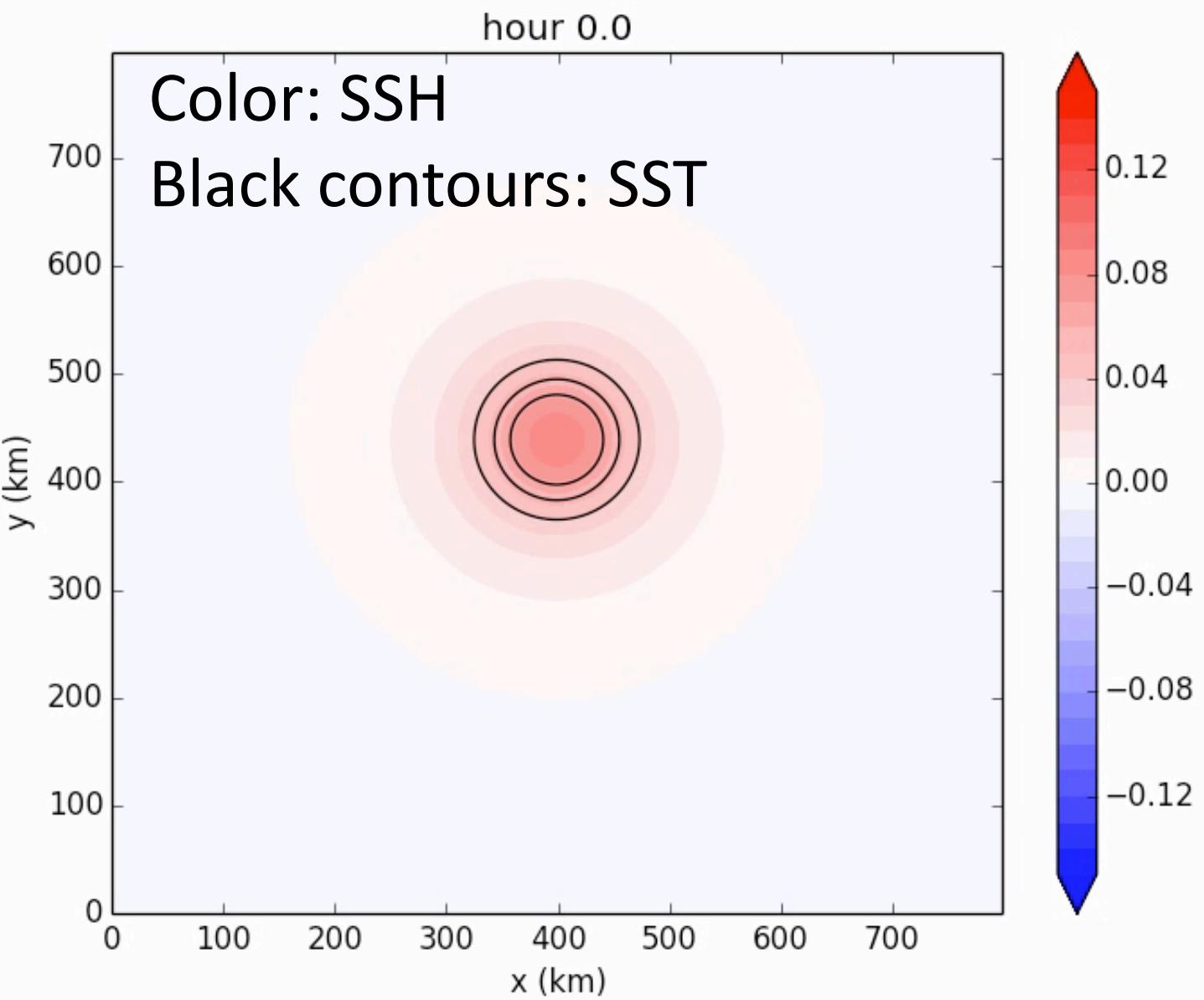
$$L_x = L_y = 800 km$$

$$R_d = 70 km$$

$$F^y = f^{-1} U_0 \sin(\omega t) F_1(z)$$

$$at y = 50 km$$

$$\rho' = \rho'_0 \exp\left(-\frac{r^2}{60^2 km^2}\right) \phi(z)$$



Setup similar to Dunphy and Lamb (2014)

# An idealized demonstration

$$\eta = \eta^{sg} + \eta^{ig} + \eta^{iw}$$

Not considered for now  
Interior QG  
SQG                          waves

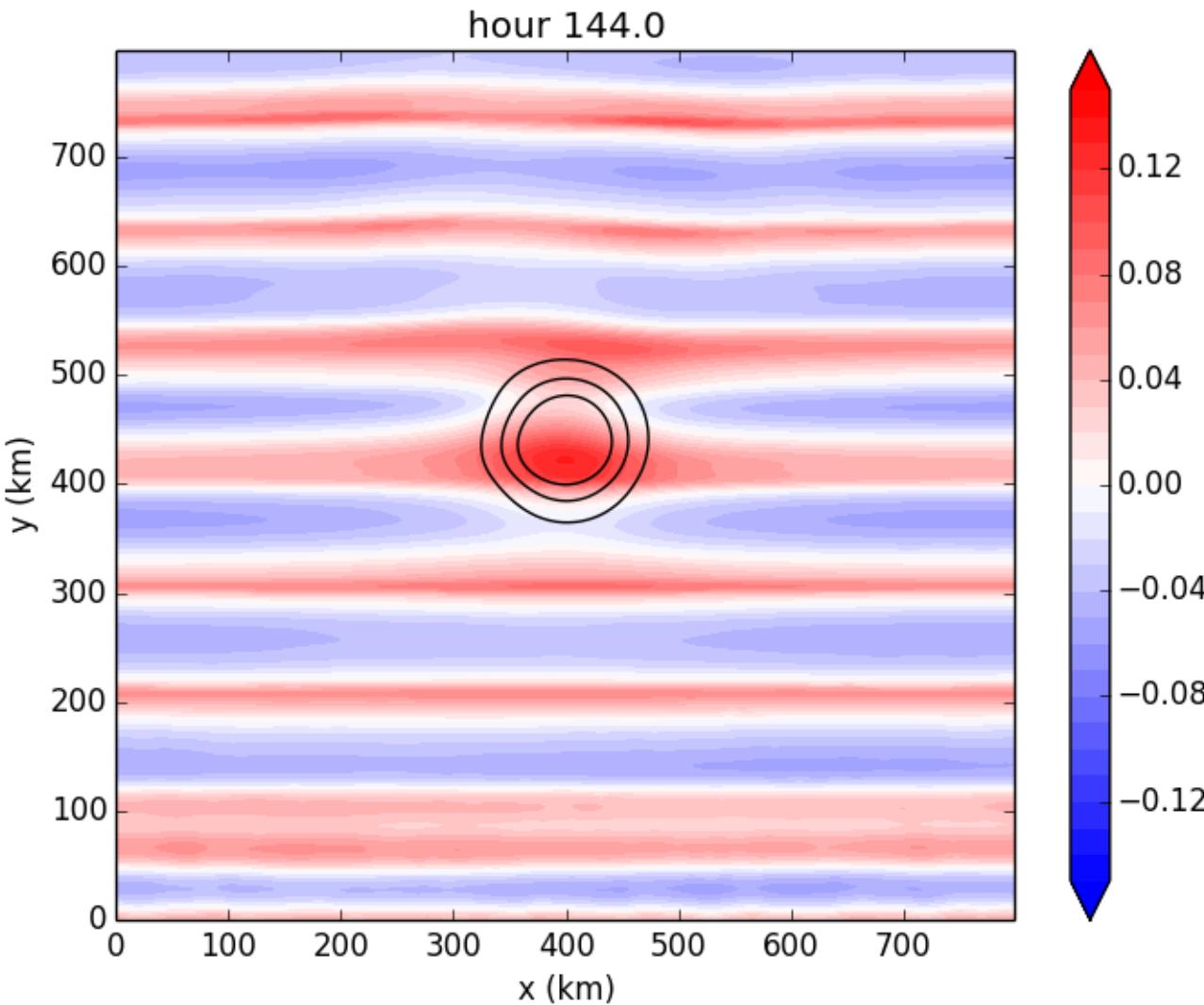
**Initial condition:**

*SSH*:  $\eta$

Density:  $\rho = \rho^{sg} + \rho^{iw}$

Horizontal velocity:

$$u_H = u_H^{sg} + u_H^{iw}$$



$\eta$  to  $\vec{u}_H$  is the tricky part

The basis of using SSH as a dynamical variable:

$$\frac{\partial u}{\partial t} - fv = -\eta \frac{\partial \eta'}{\partial x} \quad \text{SQG waves}$$
$$\frac{\partial v}{\partial t} + fu = -g \frac{\partial \eta'}{\partial y}$$

$$\eta' = \eta' [i(kx + ly - \omega t + \theta_0)]$$

$$u = \frac{k\omega + ilf}{\omega^2 - f^2} g\eta'$$
$$v = \frac{l\omega - ikf}{\omega^2 - f^2} g\eta'$$

$$\omega \ll f: \text{geostrophic approximation: } u = -\frac{g\eta_y}{f}, v = \frac{g\eta_x}{f}$$

$\omega \geq f: \text{inertia - gravity waves}$

Gill 1982, pp.257-263

# $\eta$ to $\vec{u}_H$ is the tricky part

The basis of using SSH as a dynamical variable:

$$\begin{aligned}\frac{\partial u}{\partial t} - fv &= -g \frac{\partial \eta^{iw}}{\partial x} \\ \frac{\partial v}{\partial t} + fu &= -g \frac{\partial \eta^{iw}}{\partial y} \quad \text{unknown} \\ \eta^{iw} &= \hat{\eta}^{iw} [i(kx + ly - \omega t + \theta_0)]\end{aligned}$$

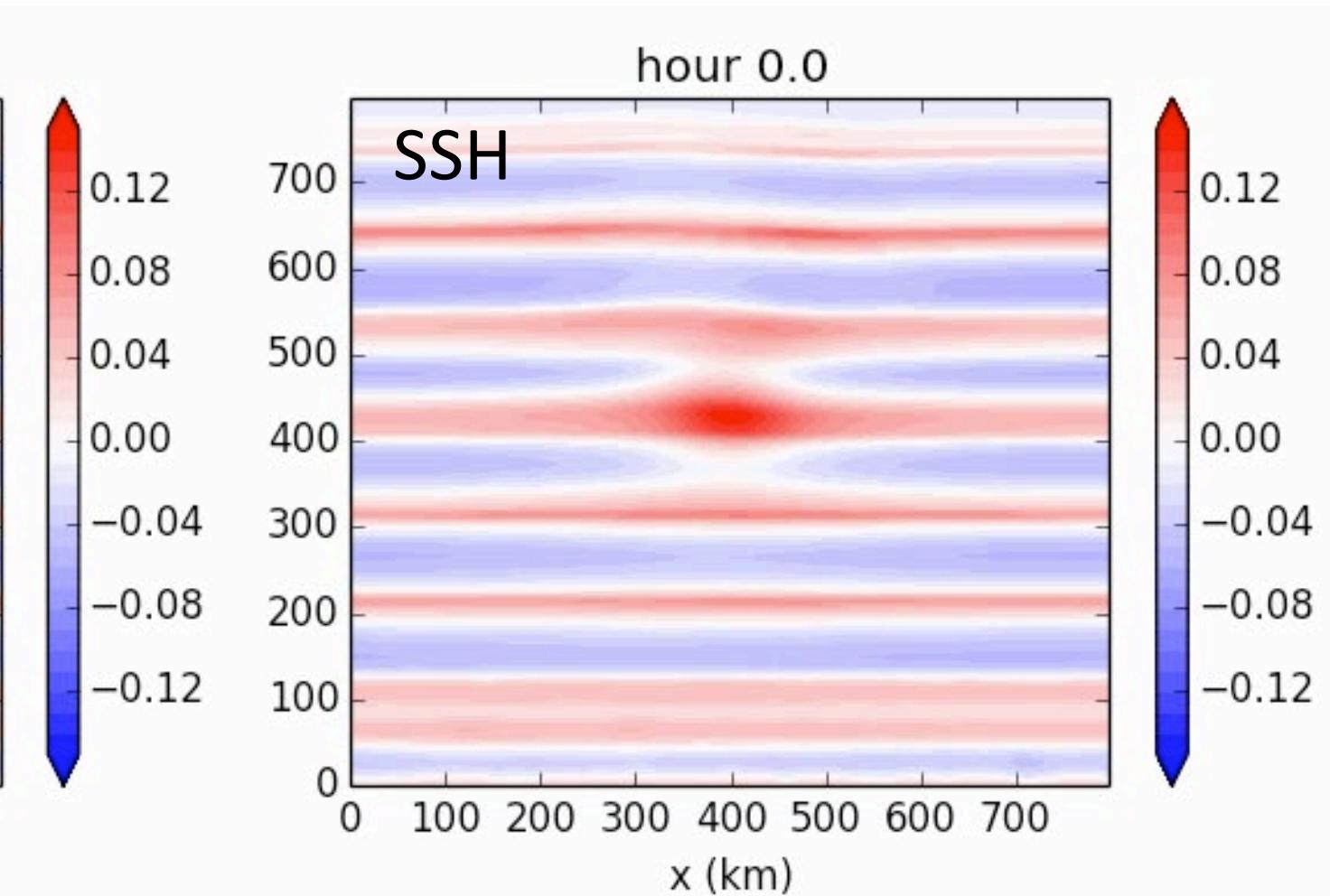
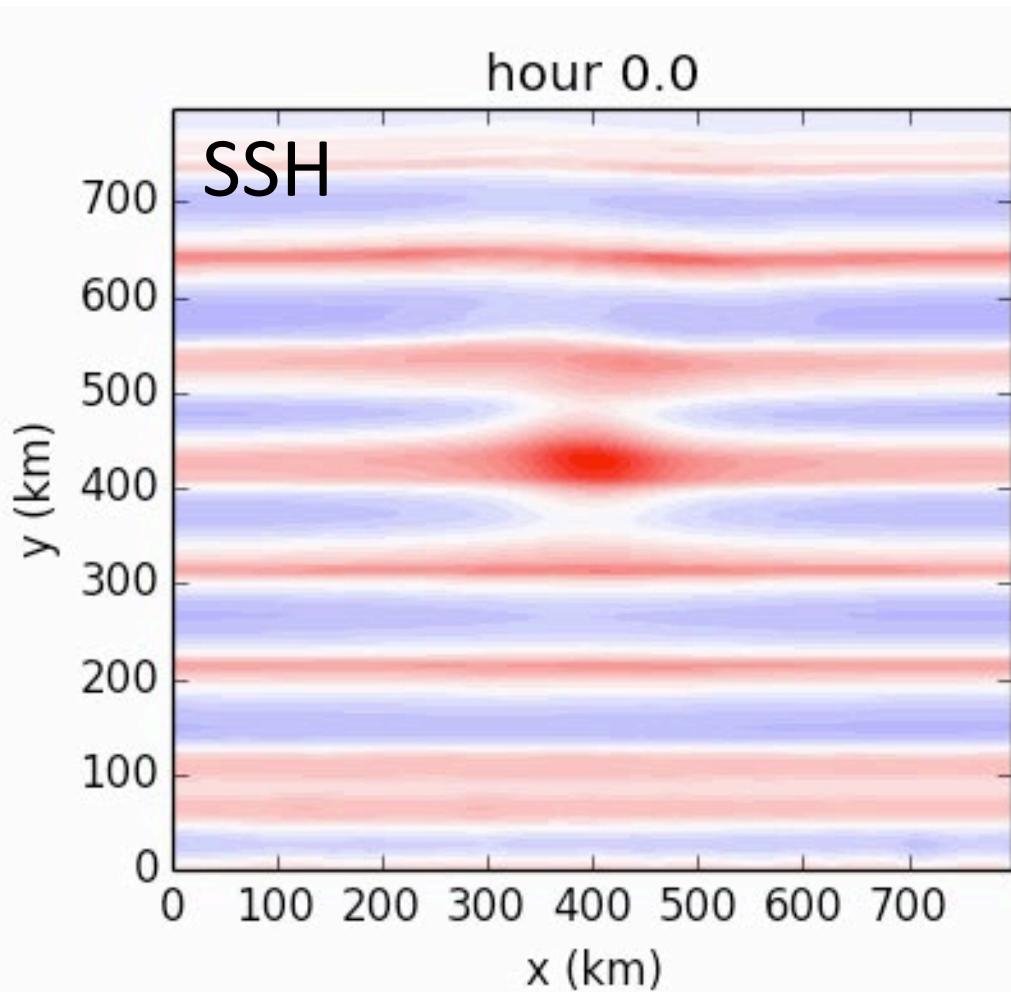
$$\begin{aligned}\hat{u} &= \frac{k\omega + ilf}{\omega^2 - f^2} g \hat{\eta}^{iw} \\ \hat{v} &= \frac{l\omega - ikf}{\omega^2 - f^2} g \hat{\eta}^{iw}\end{aligned}$$

$$\omega \ll f: \text{geostrophic approximation: } u^{sg} = -\frac{g\eta_y^{sg}}{f}, v^{sg} = \frac{g\eta_x^{sg}}{f}$$

$\omega \geq f$ : inertia – gravity waves

Gill 1982, pp.257-263

$$\eta; \rho = \rho^{sg} + \rho^{iw}; u_H = u_H^{sg} + u_H^{iw}$$



Cumulative time average since hour 0

# $\eta$ to $\vec{u}_H$ is the tricky part

The basis of using SSH as a dynamical variable:

$$\frac{\partial u}{\partial \nu} = -g \frac{\partial \eta^{iw}}{\partial y} \quad k\omega + ilf$$

**inaccuracy in velocity reconstruction**

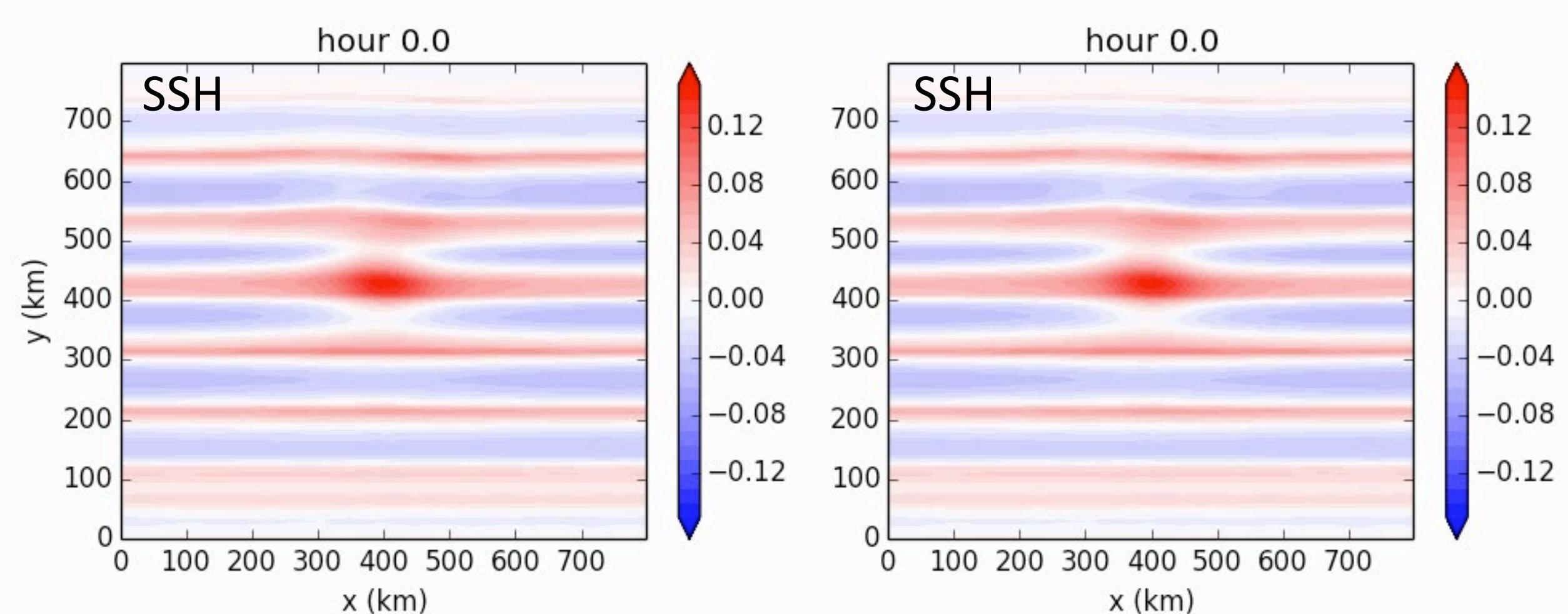
$$\frac{\partial u}{\partial \nu} + fu = -g \frac{\partial \eta}{\partial y}$$
$$\eta^{iw} = \hat{\eta}^{iw} [i(kx + ly - \omega t + \theta_0)]$$
$$\hat{v} = \frac{l\omega - ikf}{\omega^2 - f^2} g \hat{\eta}^{iw}$$

$$\omega \ll f: \text{geostrophic approximation: } u^{sg} = -\frac{g\eta_y^{sg}}{f}, v^{sg} = \frac{g\eta_x^{sg}}{f}$$

$\omega \geq f$ : inertia – gravity waves

Gill 1982, pp.257-263

$$\eta; \rho = \rho^{sg} + \rho^{iw}; u_H = \underline{u}_H^{sg} + \underline{u}_H^{iw} u_H(z=0) F_1(z)$$



Cumulative time average since hour 0

# Conclusions and discussions

$$\eta = \eta^{sg} + \eta^{ig} + \eta^{iw}$$

- SSH, SST, sea surface velocity all together can produce accurate initial condition for regenerating the time evolution of SSH.
- High resolution SST measurements are not always accessible. More interior data such as Argo measurements will help better reconstruct the vertical structure.
- It is deterministic in this study but an optimization problem in reality. The framework can be used to design a simplified dynamic core for high-resolution data assimilation.
- Is it possible to distinguish  $\eta^{ig}$  and  $\eta^{iw}$  from SSH snapshots? (we need new insight).

# One more step towards the reality

work in progress

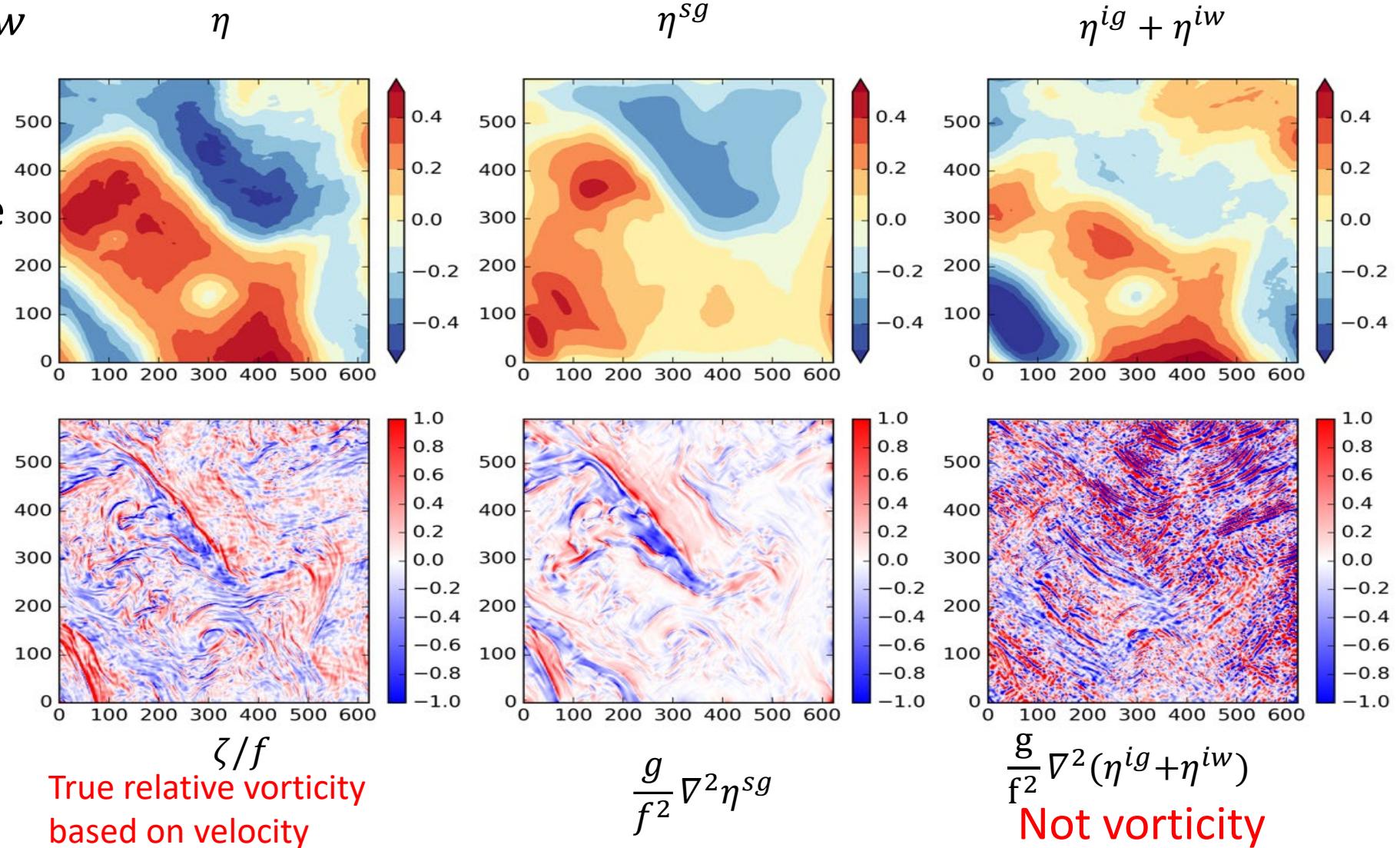
$$\eta = \eta^{sq} + \eta^{ig} + \eta^{iw}$$

Based on  $\frac{1}{48}^\circ$  MITgcm

Interior balanced mode  
is mixed in  $\eta - \eta^{sq}$

$$\frac{\partial \zeta}{\partial t} = f \frac{\partial w}{\partial z}$$

Relative vorticity is  
influenced by internal  
waves through  
stretching



# Challenge 1. Gap filling

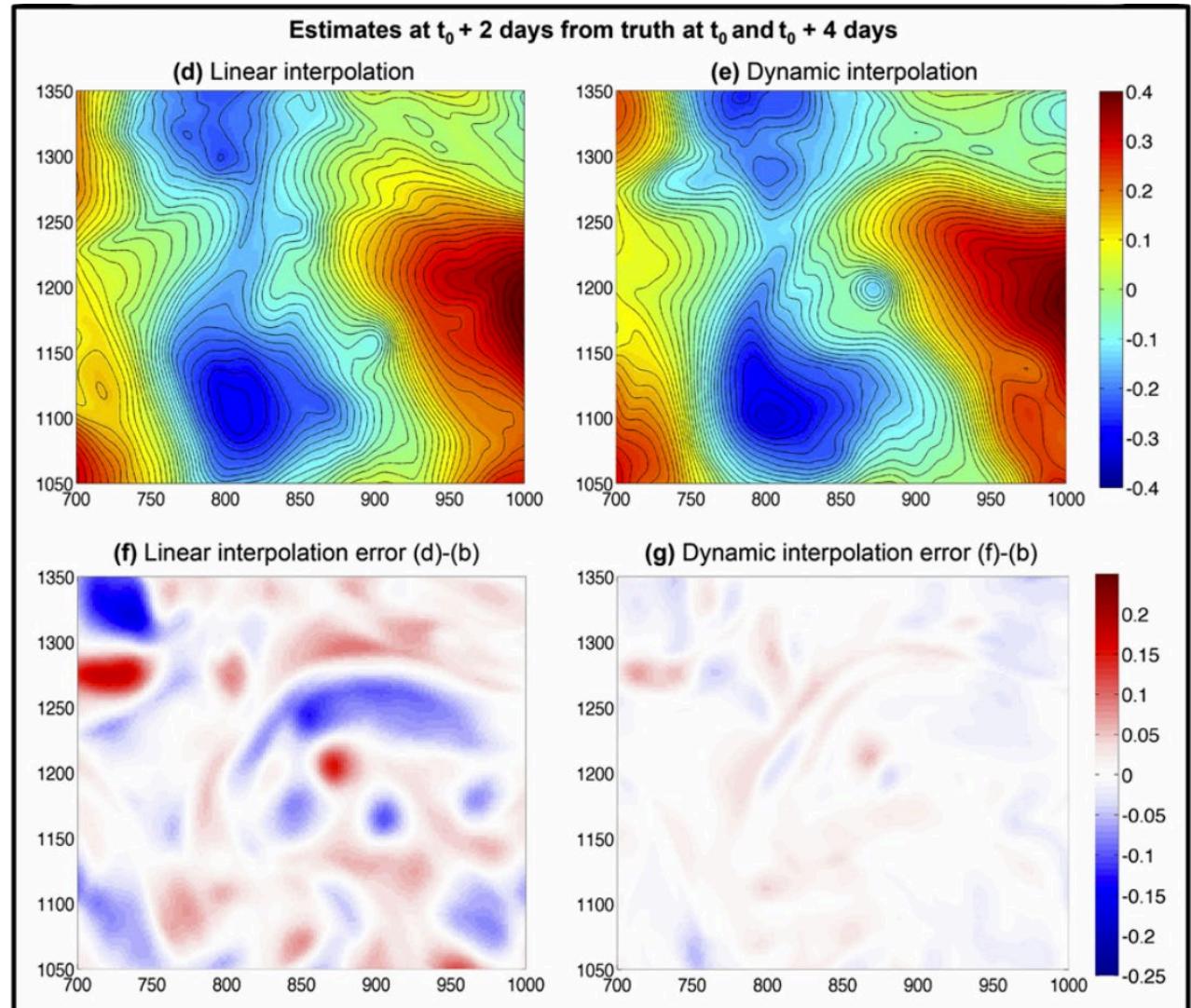
Dynamic interpolation

(Ubelmann et al. 2015,2016)

Use the 1½ layer QG model to horizontally advect information to fill the temporal gaps (PV dynamics)

**Remaining challenges:**

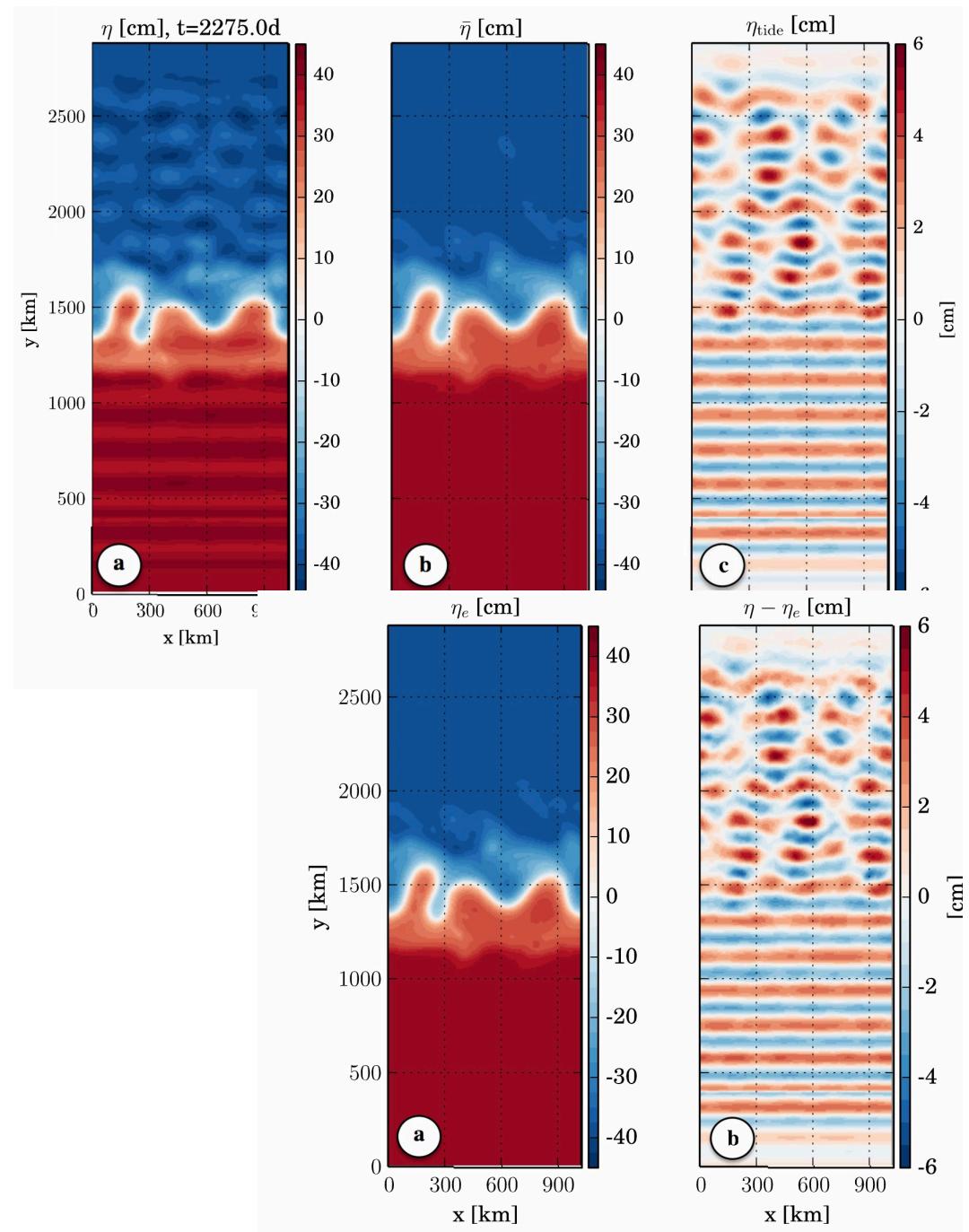
1. Account for the surface PV contribution (SQG)
2. Capture the multi-modal physics
3. Include unbalanced high frequency internal waves



# Wave-eddy separation from SSH snapshots

Ponte et al. (2017)

1. Surface density for balanced motion (eSQG framework)
2. The residual is due to waves

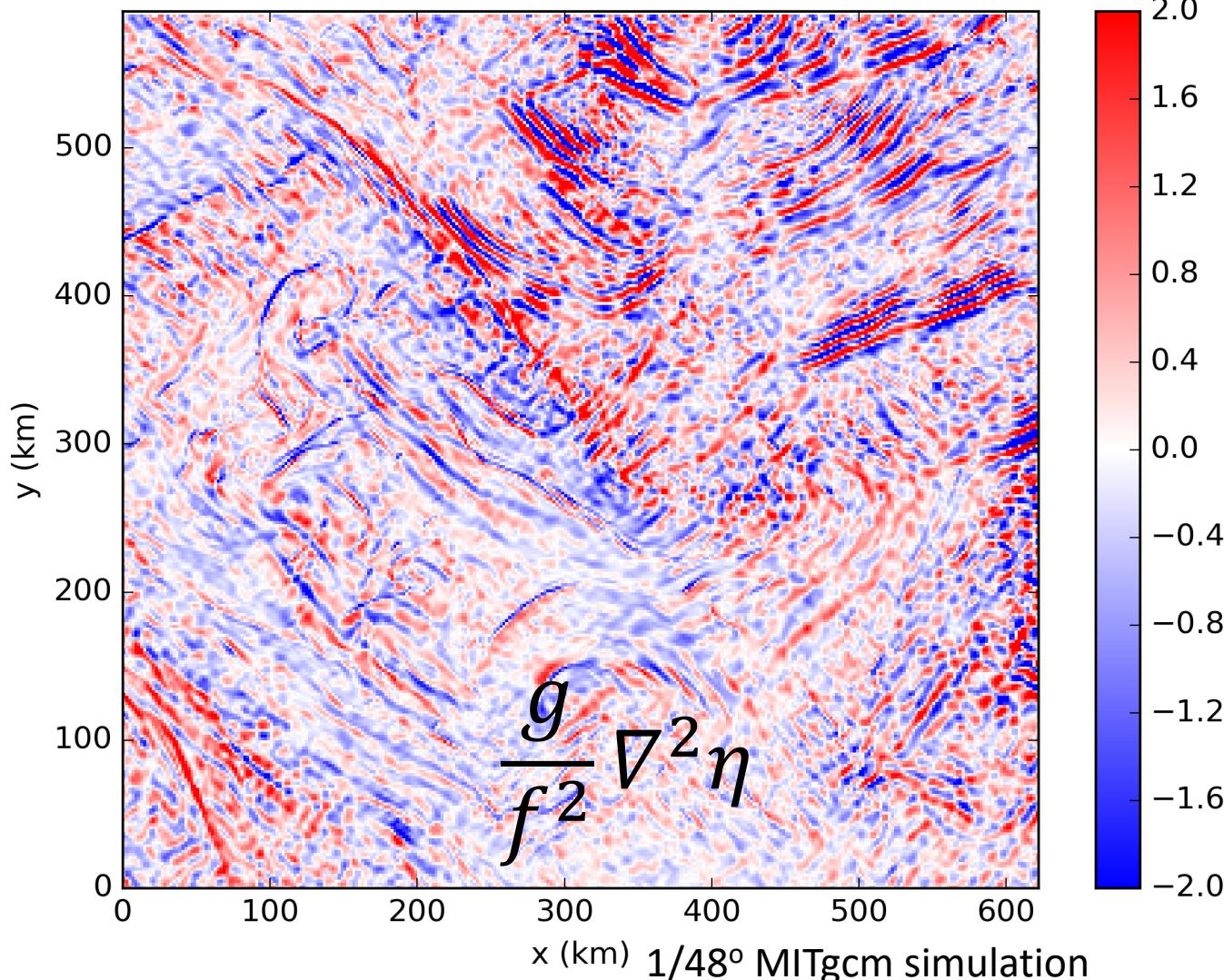


# The breakdown of geostrophic balance in SSH due to high frequency waves

$$\nabla^2 \psi \neq \zeta$$

The simple Laplacian of  
SSH is not relative vorticity.

pressure potential  
contributes to the velocity  
tendency



# Reconstruction in the two dynamical regimes

$\omega \ll f$ , rotational component

QG dynamics

$$\frac{dq}{dt} = 0$$

$$q = \nabla^2 \psi + \frac{f^2}{N^2} \partial_{zz} \psi$$

$$u = -\psi_y, v = \psi_x$$

$$\partial_z \psi = \frac{\rho'}{f}$$

$$\nabla^2 \psi + \frac{f^2}{N^2} \partial_{zz} \psi = Q$$

homogeneous solution (interior mode)  
+ particular solution (SQG)

$\omega \geq f$ , propagating component

Wave dynamics

$$(\partial_t^2 + f^2) \partial_z^2 w + [N^2(z) + \partial_t^2] (\partial_x^2 + \partial_y^2) w = 0$$

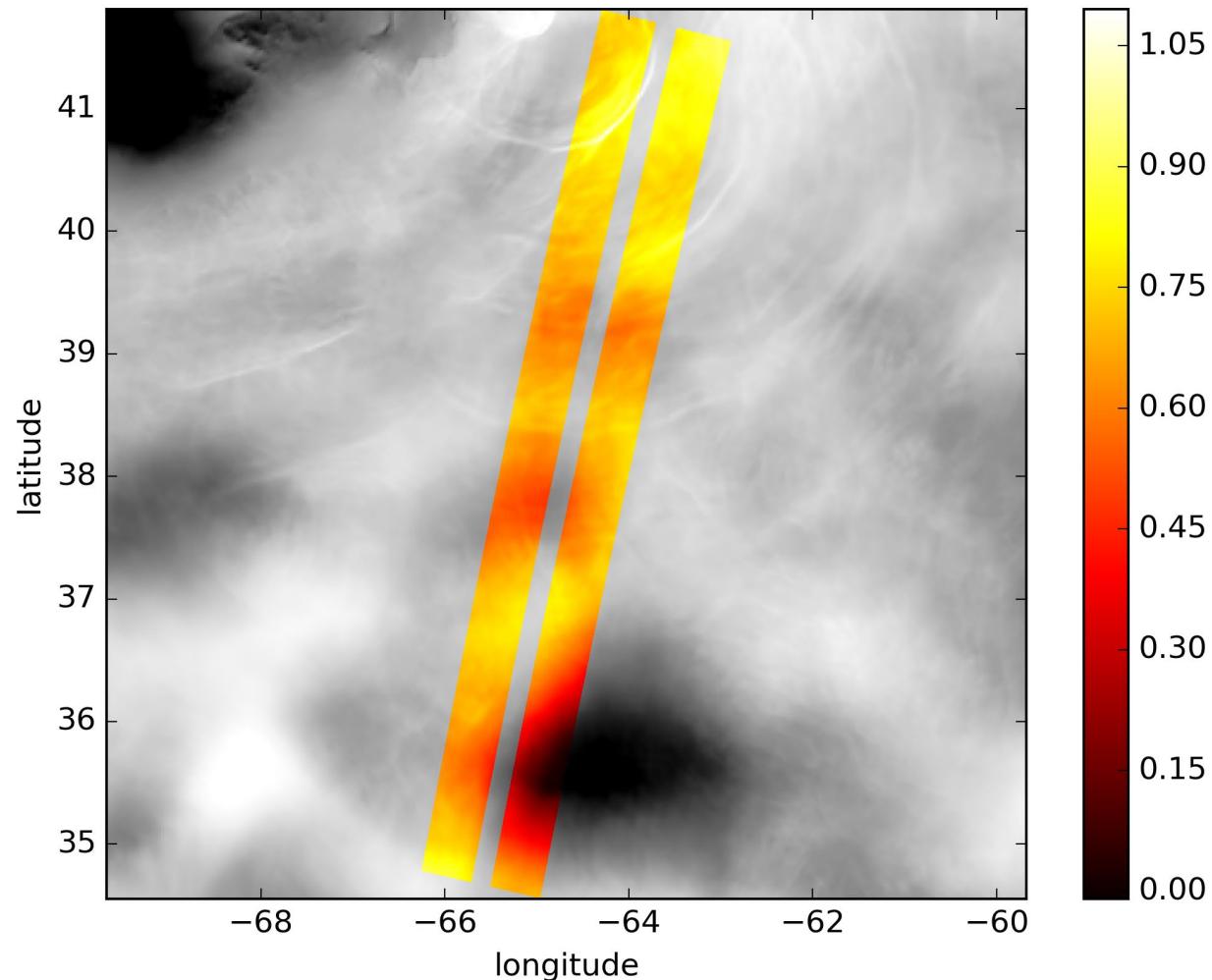
Polarization relations with a constant N

$$u = \frac{k\omega + ilf}{\omega^2 - f^2} \frac{p'}{\rho_0} \quad v = \frac{l\omega - ikf}{\omega^2 - f^2} \frac{p'}{\rho_0}$$

$$\rho' = \rho_0 \frac{-imN^2}{g(N^2 - \omega^2)} \frac{p'}{\rho_0}$$

$$\frac{\partial}{\partial z} \left( \frac{f_0^2}{N^2} \frac{\partial F_m}{\partial z} \right) = -R_m^{-2} F_m; \quad \frac{dF_m}{dz} = 0 \quad \text{at} \quad z = 0, -H.$$

# A challenge in using SWOT SSH



A snapshot of SSH from the 1/48<sup>th</sup> MITgcm

# interior+surface Quasi-geostrophy+Waves (isQGW)

Based on Wang et al. (2013) and Ponte et al. (2017)

$$\eta = \eta^{sg} + \eta^{ig} + \eta^{iw}$$

Interior QG

SQG

waves

Share the QG dynamics

Share the baroclinic  
normal mode basis

$\eta^{ig}$  is a problem

Wang et al. (2013):  $\eta^{ig} = \eta - \eta^{sg}$

Does not distinguish  $\eta^{ig}$  and  $\eta^{iw}$

Ponte et al. (2017):  
 $\eta^{iw} = \eta - \eta^{sg}(eSQG)$

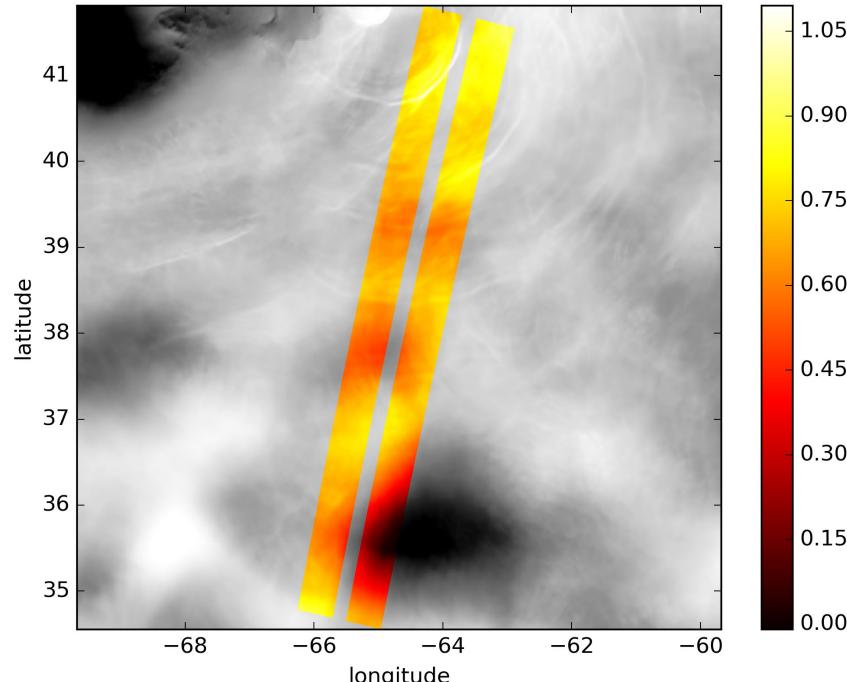
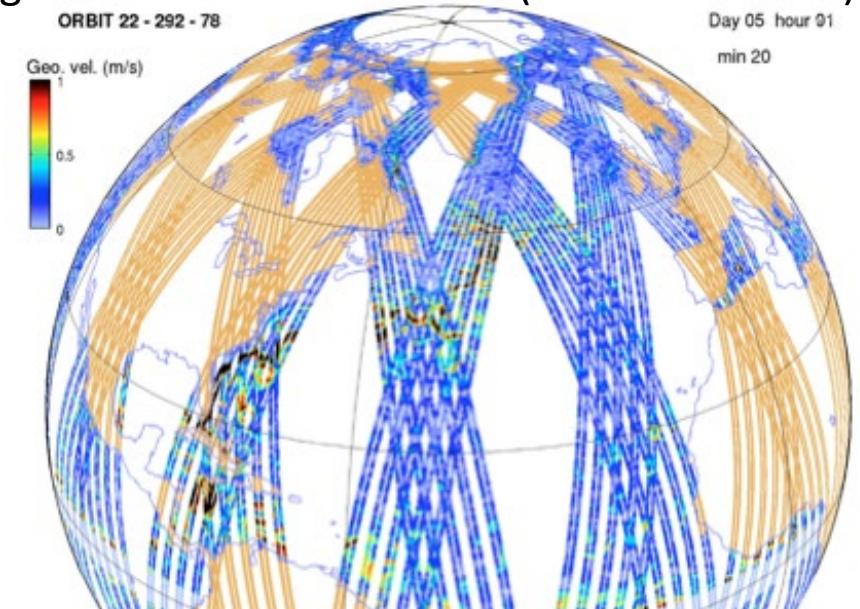
Does not consider the vertical PV tilt.

# Challenges in using SWOT SSH

1. Gaps (both in space and in time)
2. Coexistence of internal waves and low eddies coexist in SWOT SSH

## Two objectives

1. Fill the gaps
2. Separate waves and eddies from snapshot SSHs, with the aid of other available data



A snapshot of the steric height (1/48<sup>th</sup> MITgcm)