Moored time series and frequencywavenumber spectra

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SWOT will reveal an exciting new view of 10-100km SSH variability

We need a better understanding of the SSH variability at these scales. We expect internal waves to have measurable SSH signals.



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Morrow et al. (2017); modified after Rocha et al. (2016)

Salinity Processes in the Upper Ocean Regional Study

We can use mooring data and GCMs to estimate the SSH wavenumber spectrum, subject to numerous assumptions



SPURS mooring site

Domain used for analysis of JPL 2-km run of MIT-GCM



Salinity Processes in the Upper Ocean Regional Study

We had a surface mooring in SPURS measuring temperature and salinity with very good temporal and vertical resolution in the upper 400m (Farrar et al., 2015)



Frequency spectrum of surface dynamic height relative to 400m from the SPURS mooring



Dynamic height represents the dynamically driven pressure difference over some depth range

← Reference depth 500 1000 Baroclinic mode 1 1500 2000 (m) 2500 3000 3500 4000 4500 5000 -14 -12 -10 -8 -6 -2 2 -4 0 Modal pressure structure

Pressure difference between

surface and reference depth

Pressure difference between surface and reference depth (dynamic height)

Relationship of dynamic height to SSH signal of baroclinic motions



If we know how the dynamic height or SSH variance is partitioned between vertical modes, then it is easy to relate the two.



If we know how the dynamic height or SSH variance is partitioned between vertical modes, we can also convert between frequency and wavenumber using IW theory



We can convert a frequency spectrum, $\Psi(\omega)$, to a horizontal wavenumber spectrum, $\Psi(k)$, given a dispersion relation $\omega(k_h)$:

$$\Psi(k_n) = \frac{d\omega_n}{dk_n} \widetilde{\Psi}(\omega_n)$$

For internal waves:

$$\omega_n^2 = c_n^2 (k_n^2) + f^2$$

$$\frac{d\omega_n}{dk_n} = \frac{c_n^2}{\omega_n} \left(\frac{\omega_n^2 - f^2}{c_n^2}\right)^{\frac{1}{2}}$$

(c_n is a constant, different for each baroclinic mode)

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This frequencywavenumber transformation should hold for each mode, but the difficulty is to determine the relative strength of different modes

If the vertical modes are not phase locked, the SSH signal associated with mode *n* is:

 $\Psi_n^{SSH}(k_h) = a_n \Psi_n^{dyn}(k_h)$

(a_n is a known constant relating dynamic height amplitude to SSH amplitude for each baroclinic mode; $a_1=4a_2=25a_3=30a_4$) This frequencywavenumber transformation should hold for each mode, but the difficulty is to determine the relative strength of different modes

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Suppose the lowest 4 modes contribute equally to dynamic height variance (25% mode 1)



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This error bar is a factor of 3.8; if mode 1 instead contributes 95%, the estimate would be the same within error bar

Suppose the lowest 4 modes contribute equally to dynamic height (25% mode 1)

With the model, we can separate mode 1 contribution to SSH spectrum

The model and mooring estimates are roughly consistent

Conclusions

- 1. We would like to know the rough magnitude of IW signal (time-space scales of SSH) for planning SWOT cal/val and data products
- 2. The mooring estimate is most sensitive to assumed portion of dynamic height signal in mode 1 (~linearly proportional)
- 3. Both mooring and GCM estimates of SSH wavenumber spectrum are uncertain by an order of magnitude
- 4. SWOT will provide much more information on internal wave SSH than is now available

Different regions, same model (MITgcm, 2km run)

SAVAGE ET AL.: INTERNAL GRAVITY WAVE SEA SURFACE HEIGHT

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Same region, different models

Figure 9. Along-track SSH wavenumber spectra of *Jason-1* (green line) and *Jason-2* (red line) along a pass in the eastern Pacific during a time when the two satellites sampled the same ground track with a separation of 1 minute. The blue line shows the spectrum of their difference.