

# Last results and perspectives of balanced motions and internal gravity waves: **submesoscale impacts in winter**

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## Collaborators:

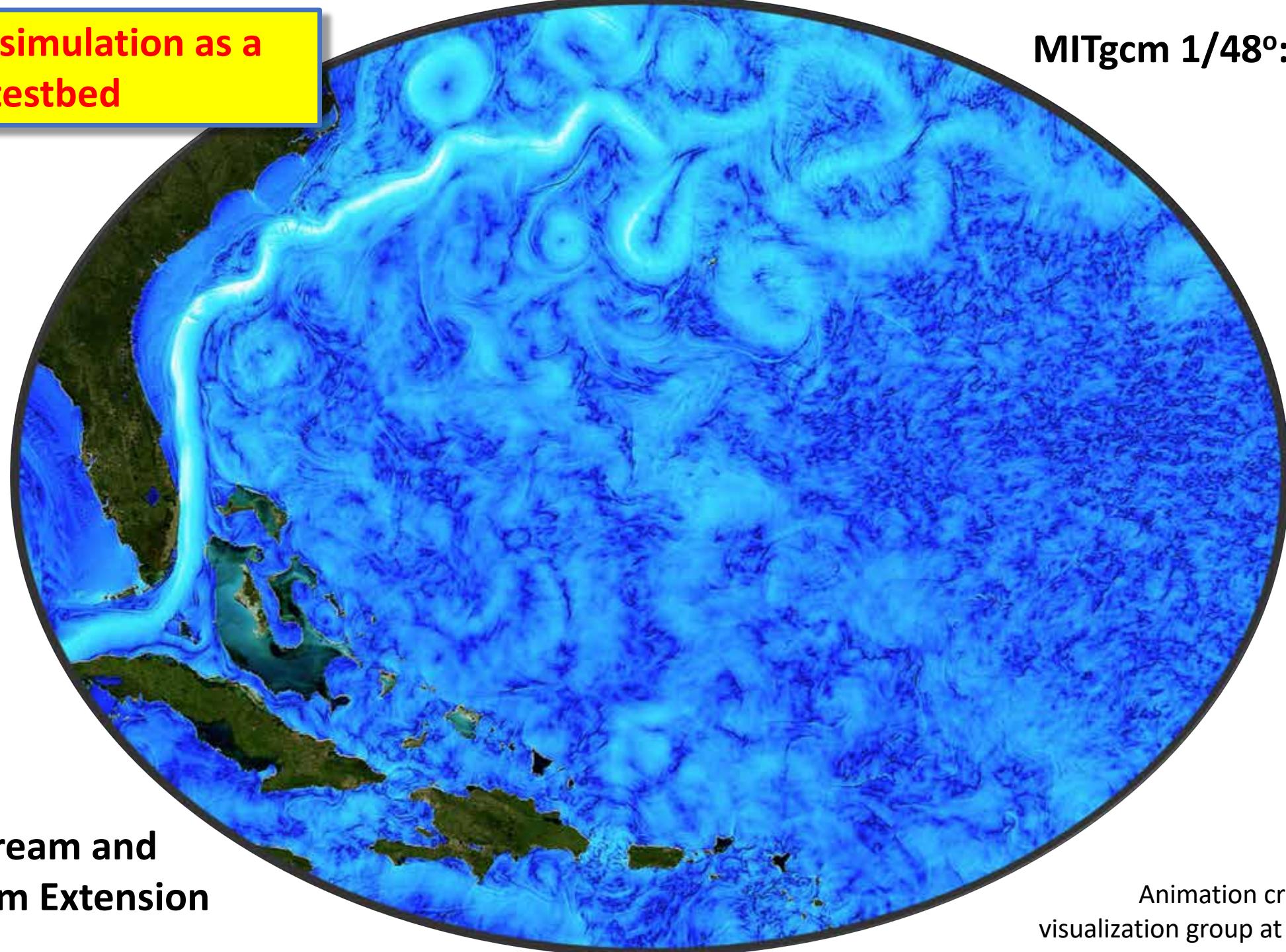
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# Outline

- 1. Helmholtz decomposition**
- 2. Surface kinetic energy during the wintertime**
- 3. Vertical heat fluxes and submesoscale motions**

**LLC4320 simulation as a  
testbed**

**MITgcm 1/48°: Surface  
speed**



**Gulf Stream and  
Gulf Stream Extension**

Animation created by the  
visualization group at NASA/AMES

# Helmholtz decomposition

$$u = -\psi_y + \varphi_x, \quad v = \psi_x + \varphi_y$$

with  $\psi$  the stream function and  $\varphi$  the potential. This leads to

$$RV = \zeta = v_x - u_y = \Delta\psi \quad DIV = -w_z = u_x + v_y = \Delta\varphi$$

In the  $\omega-k$  spectral space, the relation between KE, RV and DIV is:

$$\widehat{KE} = |\widehat{RV}|^2/k^2 + |\widehat{DIV}|^2/k^2$$

$|\widehat{RV}|^2/k^2$  and  $|\widehat{DIV}|^2/k^2$  are called respectively the RV and DIV contributions to KE.

Motivation is that this decomposition is often used to discriminate IGWs from BMS with DIV part assumed to be mostly explained by IGWs (Buhler et al. 2013; Rocha et al. 2016; Qiu et al. 2017).

The picture from the  $\omega-k$  spectra is different in winter.

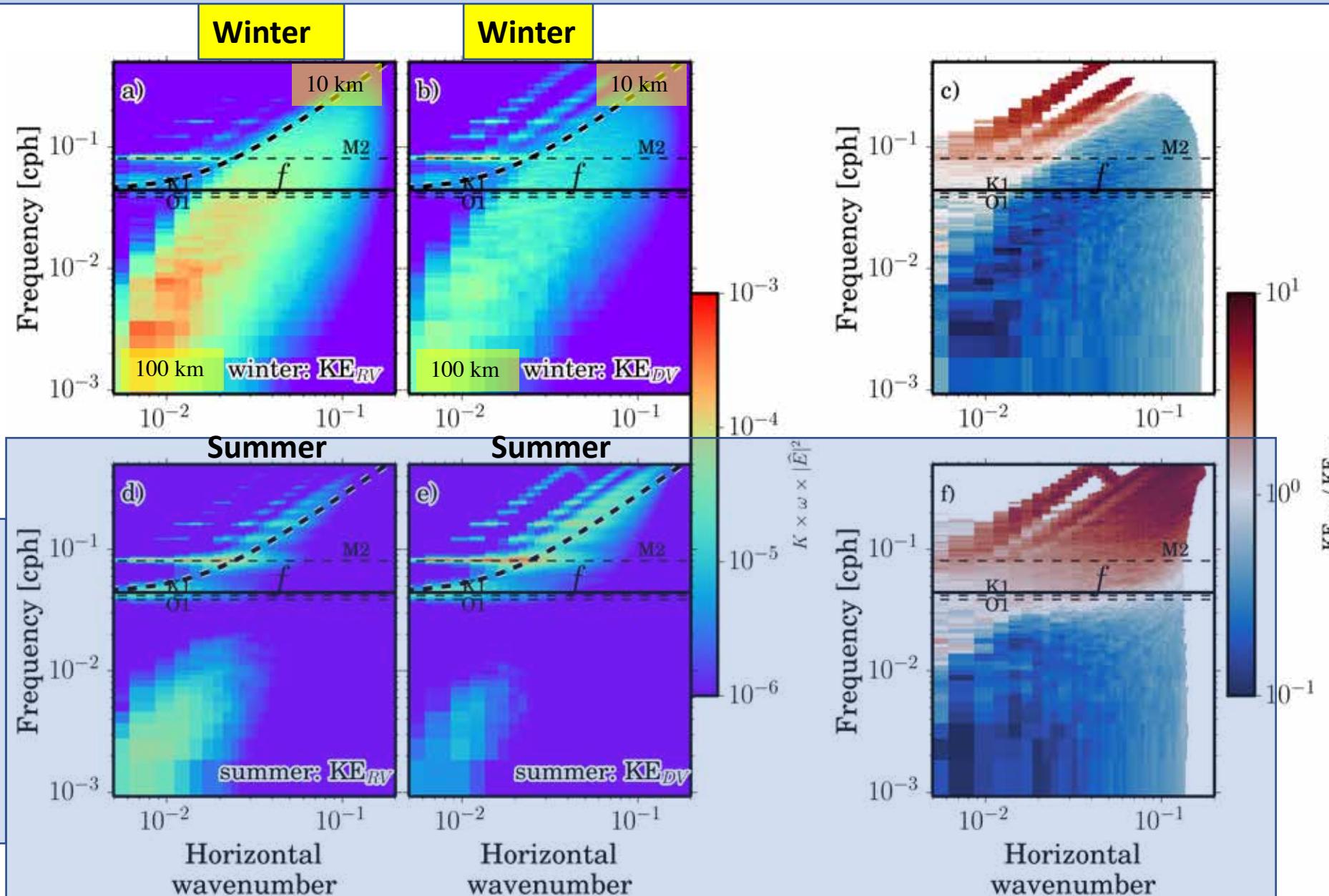
# Helmholtz decomposition: Kuroshio Extension

## RV and DIV contribution to KE

The winter RV and DIV contributions to KE are dominated by BMs

**Submesoscale  
BMs dominate  
the divergent  
part**

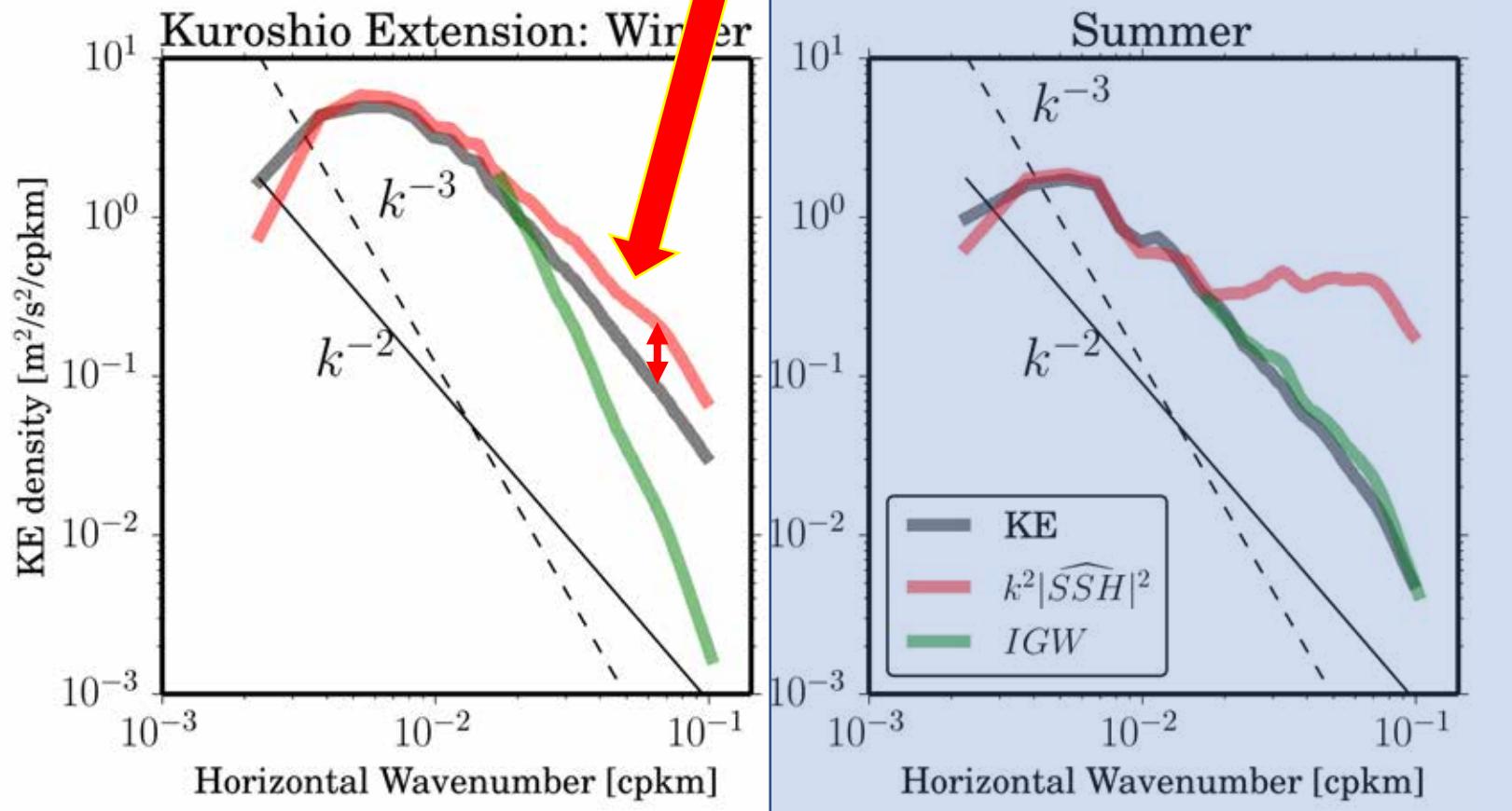
- The summer RV and DIV contributions to KE are dominated by IGWs
  - IGWs dominate the divergent part**



# Kinetic energy during the wintertime

Now, lets talk about the wintertime and the differences between  $KE_{geo}$  and KE?

➤ Geostrophic KE is larger than KE at higher wavenumbers? (not due to IGWs!)



# This is due to gradient-wind balance

$$\vec{u} \cdot \nabla \vec{u} + f \hat{k} \times \vec{u} = -g \nabla \eta$$

# cyclogeostrophic motions

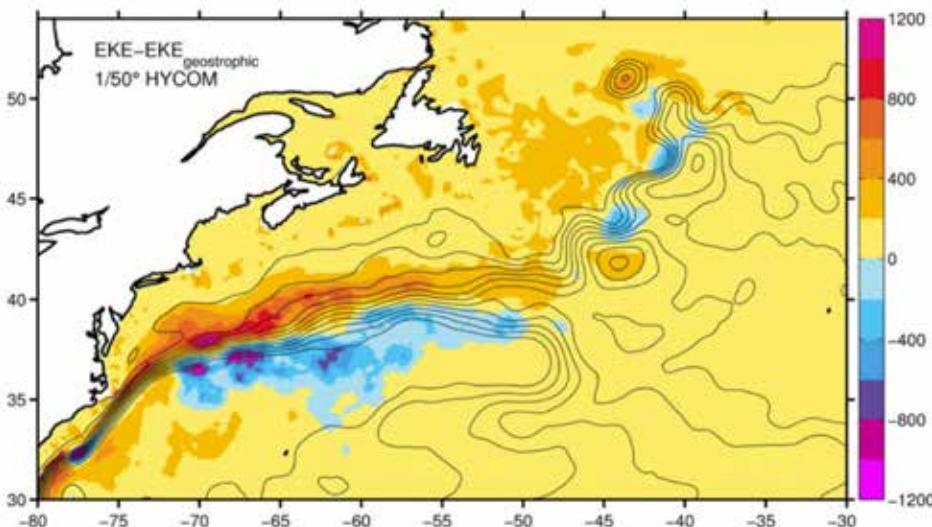
# [O(Ro)]

## Coriolis term

[O(1)]

## pressure term

[O(1)]



$$\Delta = \text{EKE} - \text{EKE}_{\text{geo}}$$

- $\Delta > 0 \rightarrow$  Cyclonic side
  - $\Delta < 0 \rightarrow$  Anticyclonic side

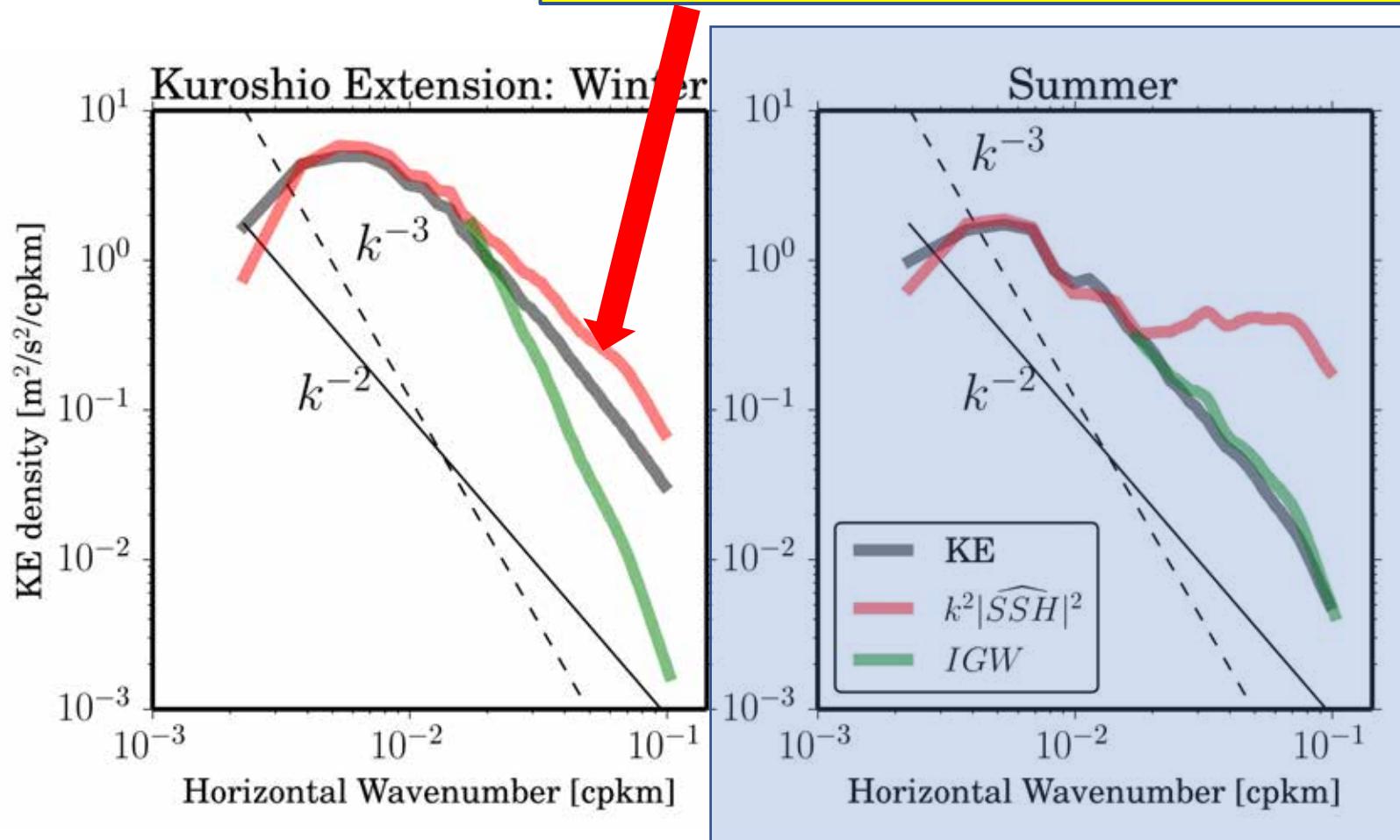
**Chassignet and Xu, 2017**

# The geostrophic balance overestimates the velocity in cyclonic motions

# Kinetic energy during the wintertime

Now, lets talk about the wintertime and the differences between SSH and KE?

The differences are explained by the gradient-wind balance:  
Cyclonic motions dominate in submesoscales (because  $\text{Ro} \sim 1$ ),  
leading to  $\text{KE} < \text{KE}_{\text{geo}}$

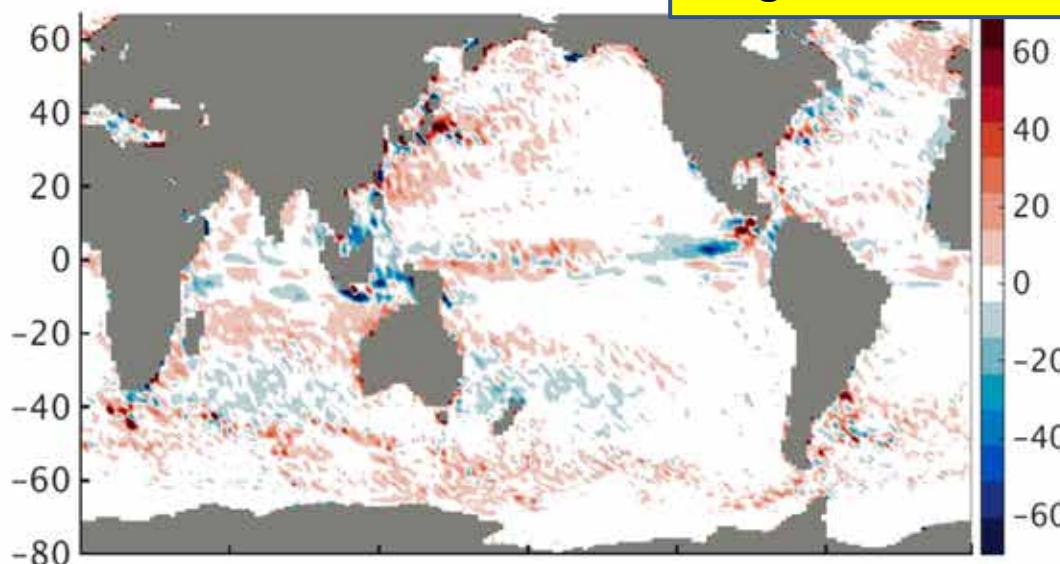
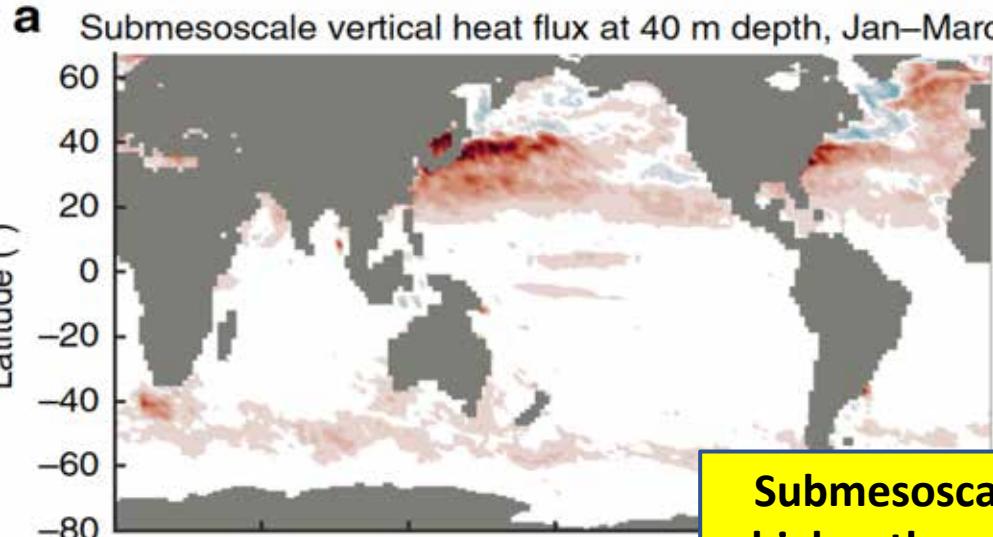


# Conclusion

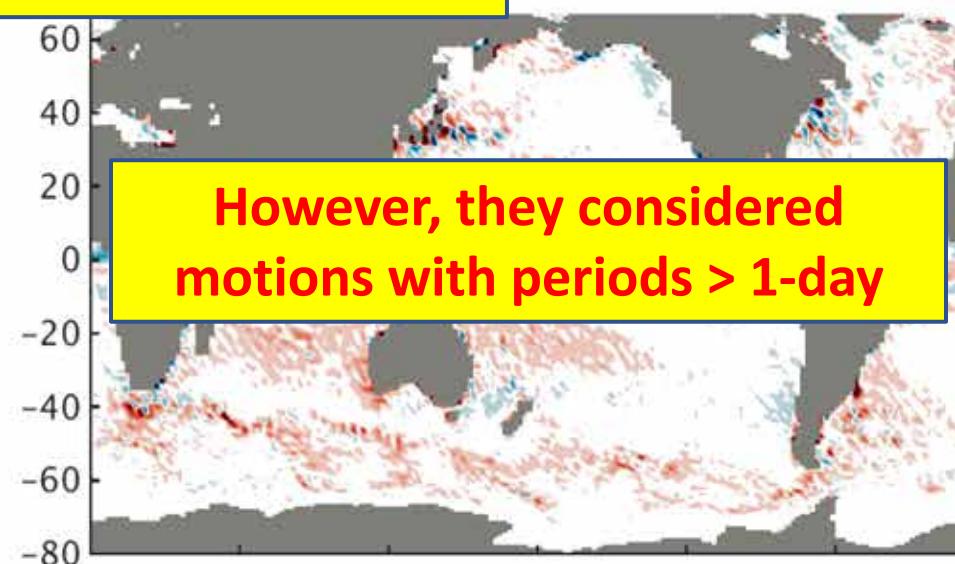
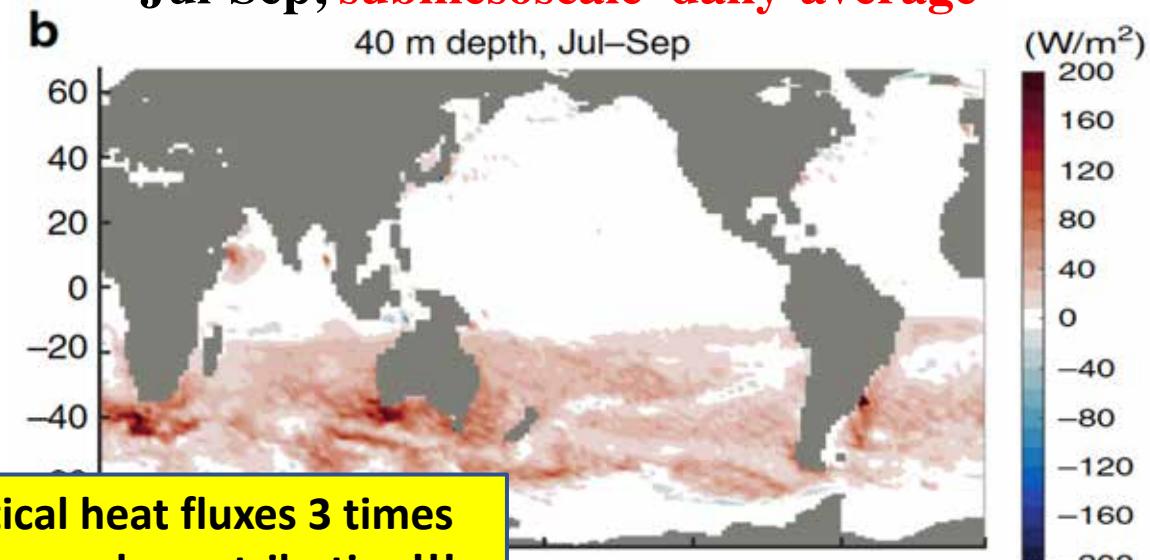
- The differences between geostrophic KE and total KE in **winter** is due to **submesoscale balanced motions in gradient-wind balance**
- The differences between geostrophic KE and total KE in **summer** is due to **internal gravity waves** (see morning's talk)

# Impact of winter submesoscales on the vertical heat fluxes?

Jan-March, submesoscale daily-average



Jul-Sep, submesoscale daily-average



However, they considered motions with periods > 1-day

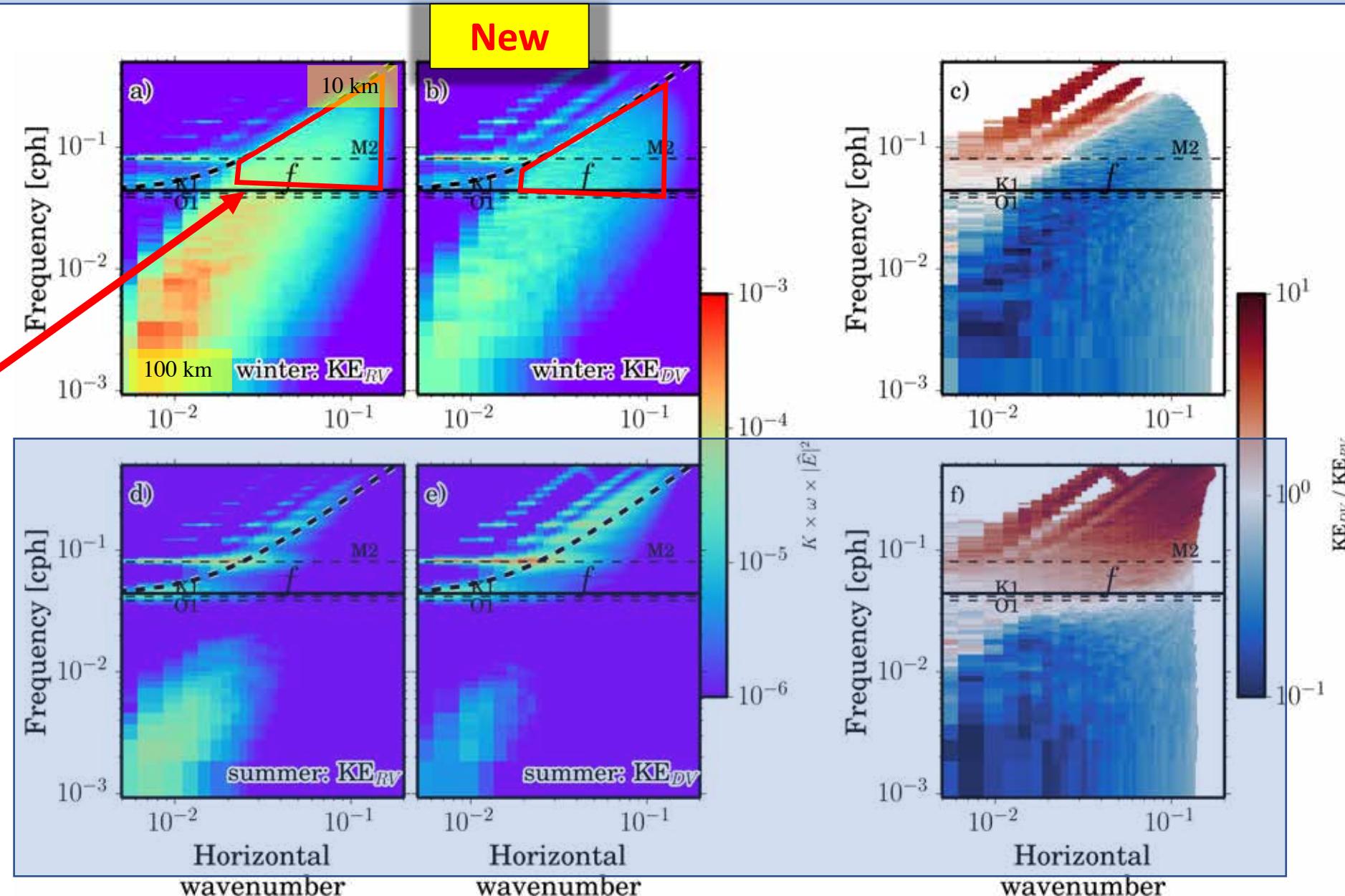
# High-frequency balanced motions

## RV and DIV contribution to KE

The winter RV and DIV contributions to KE are dominated by BMs

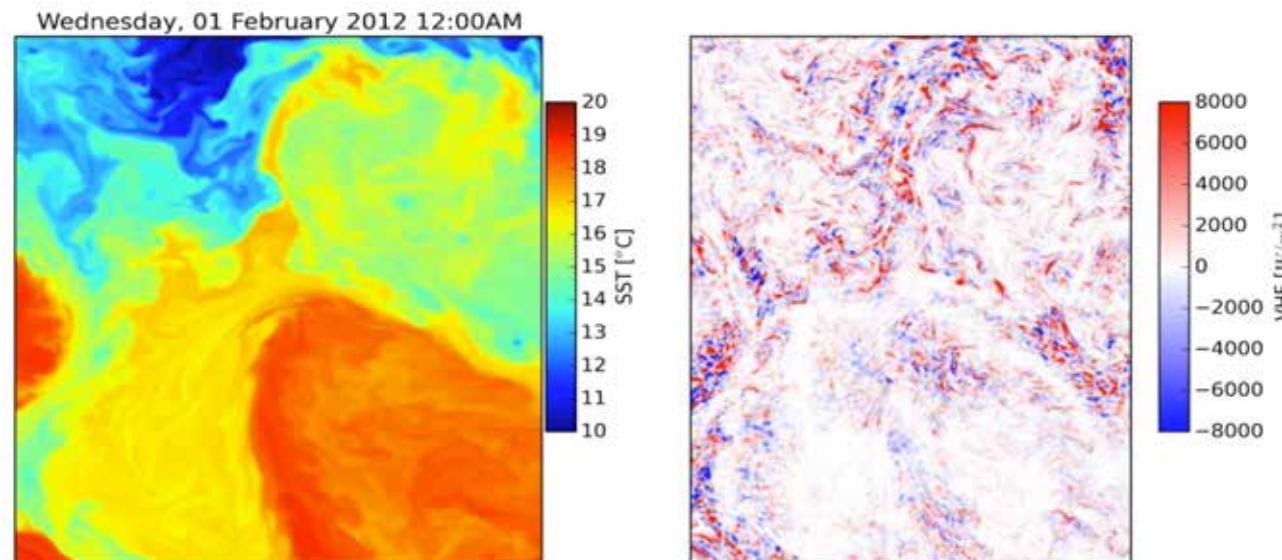
Submesoscales with frequencies higher than  $f$  are non-negligible (Thomas 2017; see Ed Zaron's talk)

- The summer RV and DIV contributions to KE are dominated by IGWs
  - The third baroclinic mode is intensified in summer



# Conclusion

- Submesoscales associated with frequencies higher than  $f$  in winter lead to almost double the vertical heat fluxes.
- Work in progress. Su et al in preparation



# Ratio divergence/rotational in the wavenumber space

- **Winter:** The spectral slope of the RV contribution in winter is  $k^2$ , as expected (Rocha et al. 2016; Qiu et al. 2017)
- **Summer:** Whereas RV contribution to KE explains most of the KE for large scales ( $> 70$  km), **DIV contribution mostly explains KE for smaller scales ( $< 50$  km)**

Ratio

$$\frac{|\widehat{\text{DIV}}|^2(k)}{|\widehat{\text{RV}}|^2(k)} = \frac{\omega^2}{f^2} = 1 + Rd^2k^2$$

Thus, the ratio increases as  $k^2$ .

